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Model simulations of inversion buildup and cold-air outflow in a small Alpine sinkhole



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Introduction

Grünloch basin

- Limestone sinkhole in the eastern Alps of Austria
- \blacktriangleright Diameter: $\approx\!\!1$ km, depth: $\approx\!\!100\text{--}200$ m
- ► Three major saddles intersect the surrounding ridgeline (Fig. 1): Lechner Saddle (≈55 m above the basin floor), Seekopfalm Saddle (≈130 m), and Ybbstaler Saddle (≈180 m).

Model simulation

Temperature inversion buildup

 Comparison of the model simulation with tethered-balloon measurements from 20 October 2008:



- An approximately 150-m deep inversion forms, with the highest stability within the lowest 50–100 m.
- The strong cooling after sunset is followed by small temperature changes later during the night.



- ▶ CM1 (Bryan and Fritsch 2002, MWR, 130, 2917–2928)
- Stretched grid: $\Delta x = \Delta y = 30-150$ m, $\Delta z = 10-400$ m
- The simulation is initialized with a quiescent and dry atmosphere.
- ► The model topography is a simplified and smoothed representation of the Grünloch topography (Fig. 1): Lechner Saddle (≈50 m above the basin floor) and Seekopfalm Saddle (≈150 m).







Fig. 2: Vertical profiles of temperature, wind speed, and wind direction in the Grünloch from tethersonde observations (left) and the model simulation (right).

Fig. 3: Time series of simulated and observed temperatures at 5 and 41 m AGL (top) and potential temperature tendency terms at 5 m AGL (bottom).

- A jet-like wind profile forms, with maximum wind speeds near the top of the layer with highest stability and near the height of the Lechner Saddle.
- The wind turns from southerly to southeasterly at the height of the jet maximum, that is, towards the direction of the Lechner Saddle.

Cold-air outflow

- Air flows out of the basin through the Lechner Saddle, mostly below the height of the Seekopfalm Saddle.
- Above the height of the Seekopfalm Saddle, the flow is mostly directed into the basin along the surrounding topography.



Fig. 4: Wind speed perpendicular to the surrounding topography. The cross section is along the blue line in Fig 1.

- The katabatic flow along the east sidewall separates from the surface near the top of the layer with highest stability.
- A jet-like wind structure over the basin connects the katabatic flow on the east sidewall and the outflow through the Lechner Saddle.



Mass budget



Fig. 6: Vertical velocity at the height of the Seekopfalm Saddle within the area indicated by the red line in Fig. 1.

- Comparison of the outflow through the Lechner Saddle with the inflow through katabatic flows along the sidewalls:
- 0.8

Katabatic flows along the sidewalls result in sinking motions.

- Upward motions occur in the layer adjacent to the katabatic-flow layer.
- Vertical velocities over the interior of the basin are weak.

Tendency terms

Pressure-gradient driven cold-air outflow:



Fig. 5: Horizontal cross section of wind speed and wind arrows approximately 10 m above the height of the Lechner Saddle (left) and vertical cross section of wind speed along the red line (right).

Suggested reading

Steinacker, R., and Co-authors, 2007: A sinkhole field experiment in the eastern Alps. Bull. Amer. Meteor. Soc., 88, 701-716.

Pospichal, B., S. Eisenbach, C. D. Whiteman, R. Steinacker, and M. Dorninger, 2003: Observations of the Cold Air Outflow from a Basin Cold Pool through a Low Pass. Extended Abstracts, ICAM 2003, 153-156.

Acknowledgments



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Fig. 7: Time series of volume flux through the Lechner Saddle below the height of the Seekopfalm Saddle, through a horizontal plane at the height of the Seekopfalm Saddle, and the vertical flux in the slope-wind layer at the height of the Seekopfalm Saddle.

The upward motions adjacent to the slope-wind layer are mostly balanced by the weak downward motions over the basin interior.



Fig. 8: Time-averaged vertical profiles of u and v tendency terms at the Lechner Saddle (LS) and the center of the Grünloch (CTR) between 2200 and 0400 CET.